

Formal proposal for the Bathonian GSSP (Middle Jurassic) in the Ravin du Bès Section (Bas-Auran, SE France)

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Key words: Bathonian GSSP, chronostratigraphy, biostratigraphy, geochronology, Ammonoidea, calcareous nannofossils

ABSTRACT

The Bathonian Global Stratotype Section and Point (GSSP) is proposed at the base of limestone bed RB071 (bed 23 in Sturani 1967) in the Ravin du Bès Section (43° 57' 38" N, 6° 18' 55" E), Bas-Auran area, “Alpes de Haute Provence” French department. The Ravin du Bès Section, as formal candidate GSSP for the base of the Bathonian Stage, satisfies most of the requirements recommended by the International Commission on Stratigraphy: 1) The exposure extends over 13 m in thickness. At the Bajocian-Bathonian transition, no vertical (bio-, ichno- or tapho-) facies changes, condensation, stratigraphic gaps or hiatuses have been recorded. Structural complexity, synsedimentary and tectonic disturbances, or important alterations by metamorphism are not relevant constraints. 2) There is a well-preserved, abundant and diverse fossil record across the boundary interval, with key markers (ammonites and

nannofossils) for worldwide correlation. The base of the Bathonian Stage and Zigzag Zone in Bas-Auran corresponds to the first occurrence level of *Gonolkites convergens* BUCKMAN, which coincides with the first occurrence of *Morphoceras parvum* WETZEL. Calcareous nannofossils, as secondary global marker, are present in all beds and allow characterizing the Bajocian-Bathonian transition. 3) Regional analyses of sequence stratigraphy and manganese chemostratigraphy are available. Spectral gamma-ray data corroborate an Early Bathonian deepening half-cycle of second order. 4) The criteria of accessibility, conservation and protection are assured by the “Réserve Naturelle Géologique de Haute Provence”. The Cabo Mondego Section (Portugal) is suggested as the Bathonian auxiliary section and point (ASSP) within this GSSP proposal.

Introduction

The present paper corresponds to part of the scientific dossier recently prepared by several specialists and sent to all members of the Bathonian Working Group (International Subcommittee on Jurassic Stratigraphy, November 2007), in order to achieve the formal ballot on the selection and proposal of the Global Stratotype Section and Point (hereafter GSSP) for the Bathonian Stage (Fernández-López et al. 2007b). The following persons are members of the Bathonian Working Group: Almérás Y. (France), Bardhan S. (India), Boderгат A.M. (France), Calomon J.H. (UK), Cresta S. (Italy), Dietl G. (Germany), Dietze V. (Germany), Enay R. (France), Fernández-López S.R. (Spain), Galács A. (Hungary), Hall R.L. (Canada), Henriques M.H. (Portugal), Hillebrandt A. von (Germany), Lanza R. (Italy), Mangold C. (France), Matyja B. (Poland), Meléndez G. (Spain), Mitta, V. (Russia), Mönnig, E. (Germany), Morton N. (France),

Page K. (UK), Pandey D.K. (India), Pavia G. (Italy), Poulsen N. (Denmark), Poulton T.P. (Canada), Riccardi A.C. (Argentina), Rogov M.A. (Russia), Sandoval J. (Spain), Schlögl J. (Slovak Republic), Schweigert G. (Germany), Seyed-Emami K. (Iran), Wierzbowski A. (Poland), Yin J.-R. (China). The results of the vote (December 2007) has been as follows: Total BtWG members = 33, YES votes = 31 (93.4%), NO votes = 1 (3.03), ABSTAIN = 1 (3.03%), NO RESPONSE = 0.

The Bathonian Working Group (BtWG) was established in 1984, during the 1st International Symposium on Jurassic Stratigraphy in Erlangen, in order to improve the geologic data relative to the Bajocian/Bathonian boundary. The Bas-Auran section was first proposed as stratotype in a presentation to the Luxembourg II (1967) Colloquium by Torrens, but not published until 1974 (Morton 1974; Torrens 1974a, b, 1987, 2002). This section was formally proposed as a candidate for the basal boundary stratotype of the Bathonian Stage by Innocenti et

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al. (1990) during the 2nd International Symposium on Jurassic Stratigraphy in Lisbon (1987). Over the following 20 years, several meetings have been organized by the Bathonian Working Group in Digne, La Palud, Budapest, Lyon and Torino. In the Bas-Auran area, the sections of Ravin du Bès, Ravin d'Auran and Ravin des Robines have been remeasured and recollected for taphonomic, sedimentologic and palaeoichnological analyses during 2006 and 2007. Reports from the Bathonian Working Group have been published by Mangold (from 1984 to 1999) and Fernández-López (from 2003 to 2007a) as mentioned in Fernández-López (2008). Over all the many years, no other candidate sections (besides Cabo Mondego) were judged worthy of consideration.

This dossier summarizes relevant results published by specialists, as well as comments and responses of the BtWG ballot 2007, in order to achieve the formal ballot on the proposal of the GSSP for the Bathonian Stage within the International Subcommission on Jurassic Stratigraphy.

Definition of the base of the Bathonian Stage (S.R. Fernández-López)

The Bathonian is the third of the four stages of the Middle Jurassic Series, above the Bajocian and below the Callovian. The name was introduced by d'Hallo (1843) and used as a stage by d'Orbigny (1850, pp. 607–608; 1852, pp. 491–492), derived from the “Bath Oolite”, in the vicinity of the city of Bath (SW England). *Zigzagiceras zigzag* (D'ORBIGNY 1846, p. 390, pl. 129, figs. 9–10; Arkell 1958, p. 177, text-fig. 60, 1–3) and *Gonolkites convergens* BUCKMAN (1925, pl. 546 A-B; Arkell 1956, pl. 18, fig. 8; pl. 19, figs. 1–2) are the index species, respectively, of the Bathonian basal zone and subzone. The Zigzag Zone was distinguished from the underlying Parkinsoni Zone by Oppel (1857, p. 579, 1862), and later assigned to the “Bath-Gruppe” (Oppel 1865, p. 309) in a discussion of the section at “Montagne de Crussol” in the Ardèche (France). The Convergens Subzone was mentioned by Maubeuge (1950, p. 4), based on the “Convergens horizon” that was used in letters by Arkell (1951–59, p.10; 1956, p. 62). The Bajocian/Bathonian boundary established between the Parkinsoni and Zigzag zones was recommended at the two congresses denominated “Colloque du Jurassique” held in Luxembourg in 1962 (Rioul 1964; Torrens 1965) and 1967 (Torrens 1974a, b). The localities of Bath (England) and the “Montagne de Crussol” (France), however, have been considered quite unsuitable for a typological definition of the Bathonian Stage, because they are condensed sections with discontinuous and lenticular beds (Torrens 1974a, b, 2002; Page 1996b). After the publication of Sturani (1967), the base of bed 23 of the Bas-Auran section, in which *Gonolkites convergens* BUCKMAN, *Parkinsonia pachypleura* BUCKMAN and *Morphoceras parvum* WETZEL first appear, was designated as the type by which to define the base for the Convergens Subzone of the Zigzag Zone and the base of the Bathonian Stage by several authors (Morton 1974; Torrens 1974a, b, 1987; Harland et al. 1982). Later, there was general agreement among

Bathonian specialists that the Bathonian Stage should start with the Standard Zigzag Zone, whose base is defined by the Convergens Subzone (Horizon 1 of Mangold 1984) followed by the Macrescens Subzone (Sturani 1967). The Parvum Subzone has been proposed by Mangold (1990) to denote the first Bathonian subzone of the Zigzag Zone in the Sub-Mediterranean Province, equivalent to the Convergens Subzone of the Northwest European Province and below the Macrescens Subzone. Analogously, due to palaeobiogeographical changes, the Dimorphitiformis Subzone has been proposed by Sandoval (1983) as the basal Bathonian subzone of the Zigzag Zone in the Mediterranean Province. Therefore, the position of the Bathonian basal boundary became justified by the base of the Northwest European Convergens Subzone of the Zigzag Zone remarkably well recorded in the Bas-Auran area, a place where both Convergens and Parvum subzones can be recognized and denote the rock bodies of the same stratigraphic interval (Fernández-López et al. 2007a; Pavia et al. 2008).

The Ravin du Bès Section (Bas-Auran area)

The Bas-Auran sections are located in southeastern France, in the French department “Alpes de Haute Provence”, in the Chaudon-Norante commune, around 25 km to the South-Southeast of Digne-les-Bains (Fig. 1). Three sections have been selected in two ravines (Fig. 2). The first, the Ravin du Bès Section (RB), is located near the farm *l'Amata* (coordinates: 43° 57' 38" N, 6° 18' 55" E, altitude 730 m). The second, the Ravin d'Auran Section (RA), is located in front of the farm *Bas-Auran* (coordinates: 43° 57' 29" N, 6° 19' 00" E, altitude 790 m). The third one, the Ravin des Robines Section (RR), is just 400

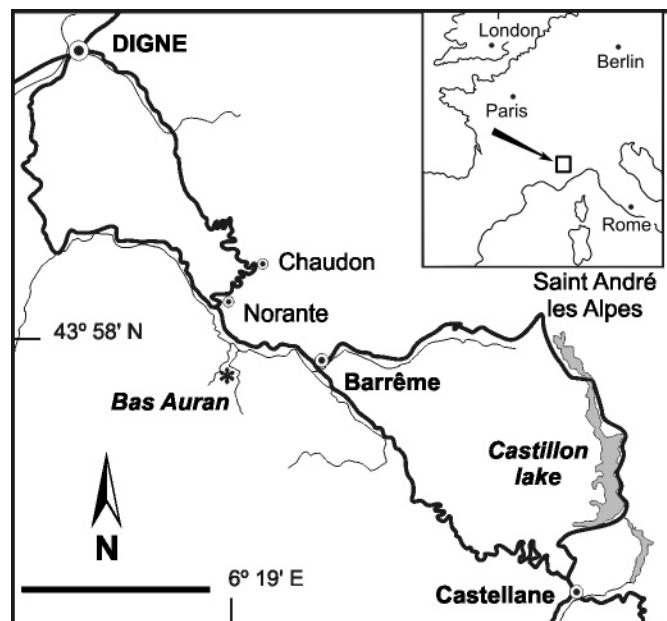


Fig. 1. Geographical setting of the Bas-Auran area (France). Scale bar is 10 km.

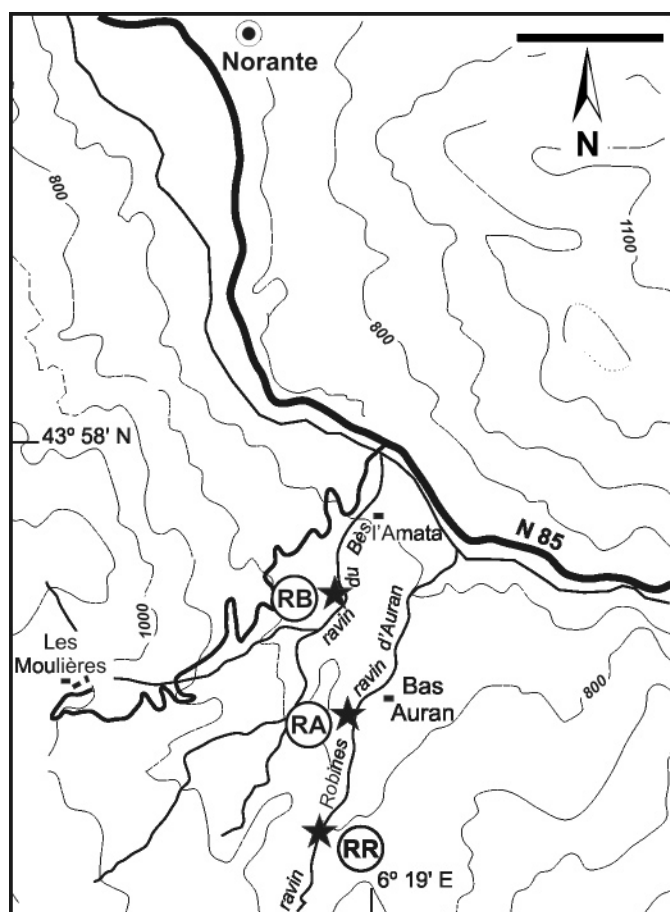


Fig. 2. Topographic sketch of the Bas-Auran area and location of the stratigraphic sections. Scale bar is 500 m. RB = Ravin du Bès Section. RA = Ravin d'Auran Section. RR = Ravin des Robins Section.

metres south of the RA section, along the Robines ravine (coordinates: 43° 57' 09" N, 6° 18' 50" E, altitude 830 m). They are located on the Castellane sheet of the "Carte géologique détaillée de la France" at the 1:80'000 scale (Goguel 1966) on the Digne sheet of the "Carte géologique de la France" at the 1:50'000 scale (Graciansky et al. 1982) and on the topographic sheet, scale 1:25'000, Barrême, no. IGN 3615.

These sections, which are free from significant unconformities, range from the Bomfordi Subzone (Parkinsoni Zone, Upper Bajocian) to the Tenuiplicatus Subzone (Aurigerus Zone, Lower Bathonian) and are over 13 m thick. Structural complexity, syngedimentary and tectonic disturbance, or considerable alterations by metamorphism, are not relevant constraints in the Bas-Auran area.

History of research on the Bathonian succession of the "Alpes de Haute Provence" and in particular of the Bas-Auran area (G. Pavia).

The area of Digne-Barrême was remarked by diverse authors as one of the most important in the world for establishing the

ammonite zonal succession of the Bathonian Stage (Garnier 1872; Haug 1891, p. 80; Guillaume 1938; Arkell 1956, p.149). The Bas-Auran locality was firstly mentioned by Haug (1891) and later visited by the French Geological Society (Zürcher 1895). Sturani in 1967 published a detailed study of the Bajocian-Bathonian succession with a litho- and biostratigraphical log settled from the whole outcrops of the Bas-Auran area. A partial revision of Sturani's work was produced by Torrens (1987) mainly on the Tenuiplicatus Subzone at the uppermost part of the marly-calcareous succession. Contribution on the lowermost Bathonian beds was finally presented by Innocenti et al. (1990) that inserted in Sturani's log the new material derived from fieldwork during ten years. More recent samplings, mainly concentrated on poorly documented and critical intervals, enlarged the Bas-Auran database from the Zigzag Zone, and furnished new and complementary results on the taphonomy of the ammonoid fossil-assemblages (Fernández-López 2007b), as well as on the taxonomy and phylogenetic arrangement of the Bathonian Bigotitinae and the origin of Zigzagiceratinae (Fernández-López et al. 2007a). Recently, Pavia et al. (2008), suggested: 1) to describe the successive ammonoid assemblages of the uppermost Bajocian to lowermost Bathonian in the Bas-Auran area, 2) to refine the subzonal biostratigraphic subdivision of the marly-calcareous succession, 3) to characterize the ammonoid content at the very base of the Zigzag Zone, 4) to demonstrate the general continuity of the ammonoid succession, and 5) to attest the suitability of one of those sections to be selected as the GSSP of the Bathonian Stage.

Geological setting of the marly-calcareous succession from the Bajocian to Bathonian in the Digne area (D. Olivero & G. Pavia)

The studied area is located in the French Subalpine Basin (FSB), corresponding to a gulf on the northwestern margin of the Tethyan Ocean (Fig. 3).

The basin is bordered by the "Massif Central" on the West and the Alpine Chain on the East (Fig. 4). During Middle Jurassic time, the basin margin was characterized by a network of tilted blocks similar to the present margin of the Atlantic Ocean (Lemoine 1984, 1985). The maximum depth of the central part of the basin probably was about 700–800 metres (Ferry 1990). The region was a transitional area between the epicontinental sea of the Paris Basin and the deep Piedmont oceanic domain, on the continental slope of the French Subalpine Basin. The studied succession is a cyclic marl-limestone alternation. In most previous works and on the geological map of Digne (Graciansky et al. 1982; Olivero & Atrops 1996) it was described as the "Calcaires à *Cancellophycus*" Formation which, in the studied region, ranges from Aalenian to Bathonian and is covered by the "Terres Noires" Formation (Late Bathonian to Oxfordian). The "Calcaires à *Cancellophycus*" Formation should not be mistaken for the partially coeval "Calcaires à *Zoophycos* du Verdon" Forma-

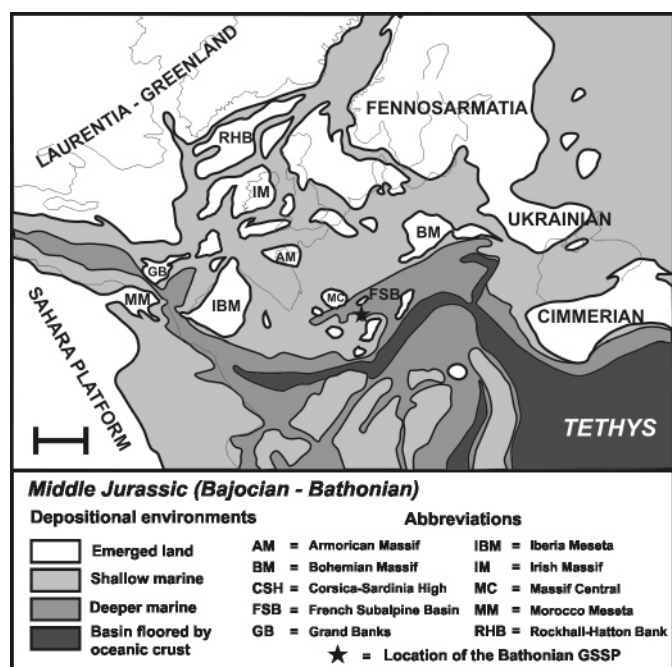


Fig. 3. The northwestern margin of the Tethyan Ocean, with the location of the French Subalpine Basin (modified from Ziegler 1988). Scale bar is 500 km.

tion, Lower Bathonian to Middle Callovian in age, proposed by Olivero and Atrops (1996) in the southernmost transitional area, between the Subalpine Basin and the Provence Platform.

Palaeoichnology, taphonomy, sedimentology and sequence stratigraphy of the upper Bajocian to lower Bathonian of the Bas-Auran area (S.R. Fernández-López & D. Olivero)

In the Bas-Auran area, Lower Bathonian deposits comprise black or grey limestone beds alternating with marls usually known as “Marno-calcaires à *Cancellophycus*” (Graciansky et al. 1982; Olivero & Atrops 1996). Petrographically and in terms of biofacies, these deposits are relatively uniform mudstones to wackestones, with common ammonoids, scarce sponges and very scarce nautiloids, brachiopods, bivalves, belemnites, echinoids, crinoids and gastropods. As to microfossils, the overall sedimentary facies shows a calcisphere-mudstone texture; the marls contain foraminifers (*Lenticulina*, *Dentalina*), ostracods and molluscs (cephalopods, bivalves, gastropods) along with detrital minerals, quartz, muscovite and biotite (Corbin et al. 2000).

Palaeoichnological studies have been carried out by Olivero (1994, 2003). Bioturbation textures are common and bioturbation structures are scarce, indicating dominant softgrounds. *Zoophycos*, *Chondrites* and *Planolites* occur from the bed RB093 to bed RB001 (Fig. 5). Local concentrations of trace fossils of these ichnotaxa in bed RB039 suggest the development of a soft- to firmground in this stratigraphic

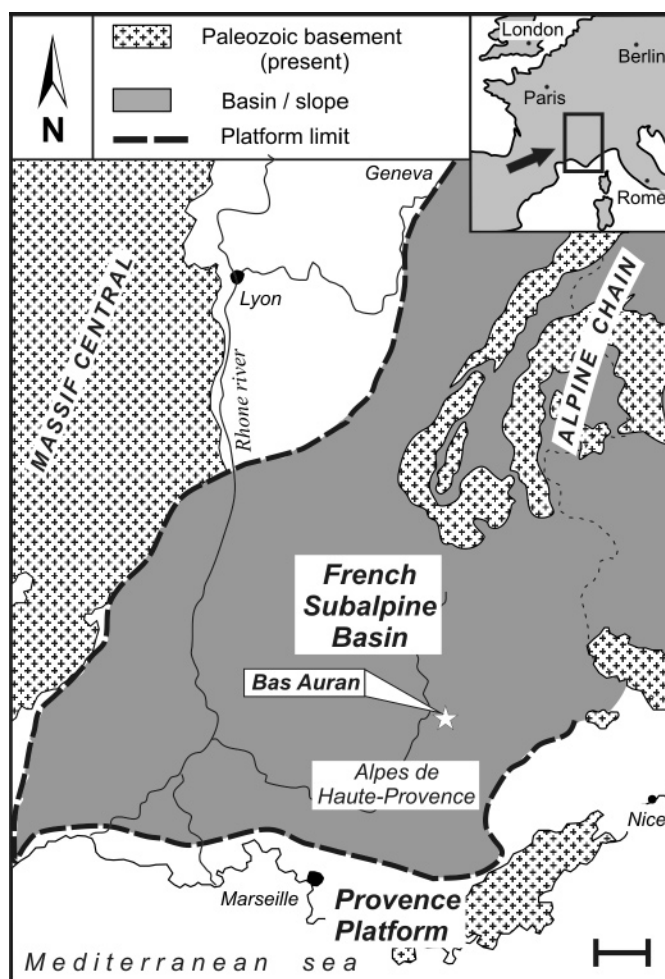


Fig. 4. The French Subalpine Basin, with the location of the Bas-Auran area (from Olivero 2003, modified). Scale bar is 25 km.

level (Fig. 5). Bioturbation structures indicative of firmground (*Thalassinoides*, *Rhizocorallium*, *Zoophycos* and trace fossils related to large *Halimedes*) occur in a more calcareous layer just overlying the top of the bed RB003. Biogenic borings indicative of hardground (*Zapfella*) are common, associated with very scarce encrusting serpulids, on the top of bed RB001, indicating the exceptional development of a stratigraphic discontinuity at the top of the “Marno-calcaires à *Cancellophycus*” on the Bas-Auran area. Sedimentation appears irregular and condensed from bed RB093 towards the top of the Bathonian Zigzag Zone, compared with previous intervals where a more constant and expanded sedimentation is suggested. At the Bajocian-Bathonian transition, however, no stratigraphic gaps or hiatuses have been recorded. From a taphonomic point of view, the occurrence of resedimented and reworked ammonoids implies that some form of current flow or winnowing affected the burial of concretionary internal moulds. Ammonoids show the following taphonomic characters at the Bajocian-Bathonian transition: 1) high values of stratigraphic persistence of

ammonoid shells, 2) dominance of homogeneous concretionary internal moulds of phragmocones, completely filled with sediment, and 3) dominance of unflattened sedimentary moulds bearing no signs of rounding, bioerosion or dense encrusting by organisms (such as serpulids, bryozoans or oysters). These taphonomic features are indicative of a low rate of sedimentation and a low rate of accumulation of sediment, associated with sedimentary starvation in deep environments (Fernández-López 2007a).

The bed-scale limestone-marl alternation is primary in origin, although accentuated by diagenetic redistribution of carbonate. Lithological differentiation between marly and limestone intervals resulted from alternating episodes of carbonate input and starvation. Both lithologies may contain evidence of sedimentary and taphonomic reworking, associated with scours, which reflect low rates of sedimentation and stratigraphic condensation. There is no evidence, however, of taphonomic condensation (i.e. mixture of fossils of different

age or different chronostratigraphic units) in the ammonoid fossil-assemblages, except in level 002. Sedimentological data and sequence-stratigraphy interpretations of the Jurassic deposits in the French Subalpine Basin have been published by Graciansky et al. (1993, 1998a, b), Olivero & Atrops 1996, Olivero et al. (1997), Hardenbol et al. (1998) and Jacquin et al. (1998).

Palaeoichnological, taphonomic and sedimentological results confirm, therefore, the development of a deepening phase associated with sedimentary starvation, within 3rd and 2nd order cycles, in the Bas-Auran area, during the Early Bathonian. The maximum deepening of a 2nd-order transgressive/regressive facies cycle (T/R 7, Upper Aalenian–Upper Bathonian, in Graciansky et al. 1993, 1998 a, b) is at the end of the Early Bathonian, which corresponds to an extensional and deepening phase of the basin. The outcrop successions at Bas-Auran show no obvious signs of non-sequence or discontinuity across the Bajocian/Bathonian boundary interval.

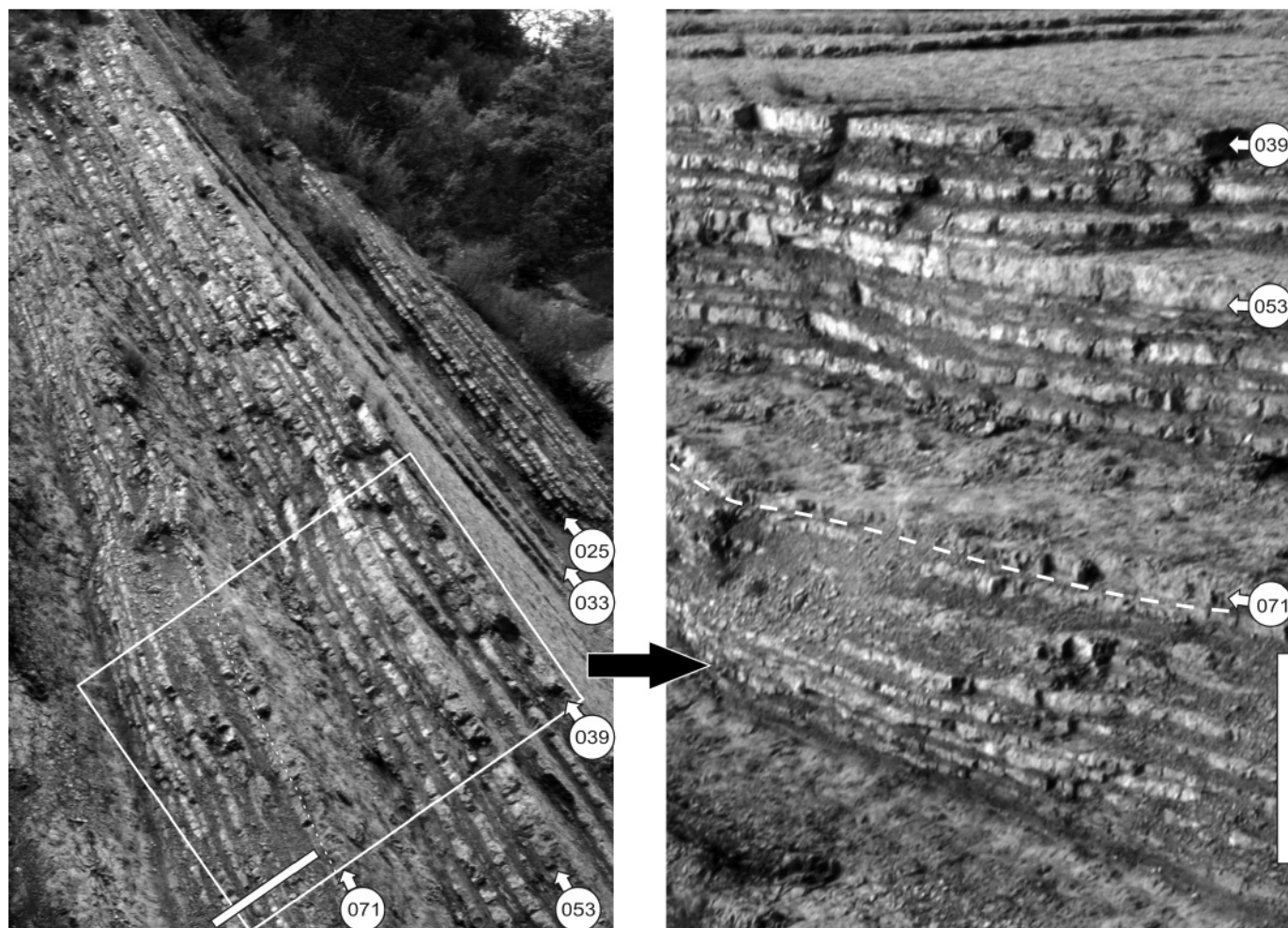


Fig. 5. Ravin du Bès Section and detail of beds around the Bajocian/Bathonian boundary interval. Limestone bed 071 indicates the base of the Bathonian. Scale bar is 1 m.

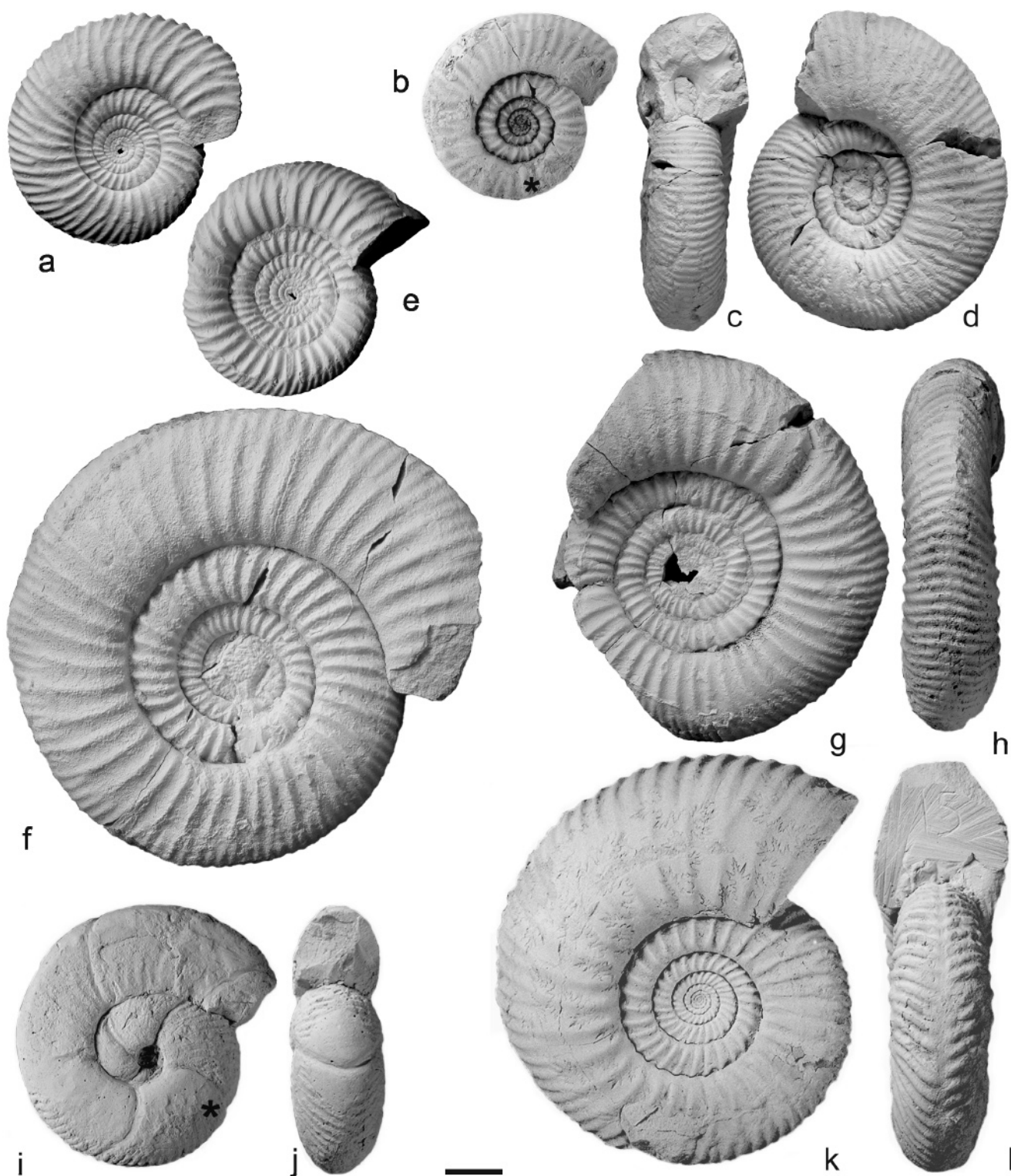


Fig. 6. Lower Bathonian ammonites from Bas-Auran area. Specimens have been whitened with magnesium oxide prior to photography. Scale bar is 1 cm. (a): *Bigotites mondegoensis* FERNÁNDEZ-LÓPEZ et al. [M], specimen PU111312, level BA14, Convergens Subzone. (b): *Protozigzagiceras* cf. *torrensi* (STURANI) [m], specimen PU111573, level RA033 (= BA12), Macrescens Subzone. Black asterisk marks the last septum of the phragmocone. (c, d): *Protozigzagiceras* aff. *torrensi* (STURANI) [M], specimen PU31694, level BA13, Convergens Subzone. (e): *Bigotites sturani* FERNÁNDEZ-LÓPEZ et al. [M], specimen PU111253, level BA19, Convergens Subzone. (f): *Bigotites diniensis* STURANI [m], specimen PU111243, level BA20, Convergens Subzone. (g, h): *Bigotites sturani* FERNÁNDEZ-LÓPEZ et al. [M], specimen PU111233, level BA20, Convergens Subzone. (i, j): *Morphoceras parvum* WETZEL [M], specimen PU111564, level BA17, Convergens Subzone. Black asterisk marks the last septum of the phragmocone. (k, l): *Gonolkites convergens* BUCKMAN [M], specimen PU111067, level BA15, Convergens Subzone.

Palaeontological records

The Bomfordi and Convergens subzones in the Bas-Auran area contain an ammonoid succession that displays a maximum value of biostratigraphic and biochronostratigraphic completeness. Additional macrofossil groups occur in the sections (e. g., sponges, bivalves, brachiopods and belemnites), although they are scarce and have not yet been studied in detail.

Ammonites (S.R. Fernández-López, C. Mangold & G. Pavia)

Biochronostratigraphic data on ammonoids of the Bas-Auran sections have been published by Sturani (1967), Pavia (1973, 1983a, b, 1984, 1994, 2000, 2007), Torrens (1987), Innocenti et al. (1990), Olivero et al. (1997) and Joly (2000). New and complementary results from the biochronostratigraphic analyses of ammonoid fossil-assemblages at the Bajocian/Bathonian boundary in Bas-Auran are in press (Fig. 6, Fernández-López et al. 2007a, Pavia et al. 2008). In the French Subalpine Basin, the successive ammonoid fossil-assemblages are composed of Mediterranean and Northwest European representatives, associated with some Sub-Mediterranean ones. Upper Bajocian and Lower Bathonian Mediterranean taxa of the suborders Phylloceratina and Lytoceratina are relatively common (up to 25% at subzonal scale, Fig. 7). Northwest European taxa, such as Parkinsoniinae, may surpass 25% at subzonal scale. Lower Bathonian Bigotitinae, endemic and characteristic of the Sub-Mediterranean Province, reach 13%. This complex palaeobiogeographical pattern of the Upper Bajocian and Lower Bathonian ammonoid fossil-assemblages enables recognition of diverse subzonal schemes and accurate chronocorrelation between the three main provinces of the West Tethyan Subrealm.

Among the possible guide fossils for the Bajocian/Bathonian boundary, Parkinsoniidae have a better record than Morphoceratidae. The lowest occurrences of *Gonolkites* [M] and *Morphoceras* [M] may be evidence of palaeobiological events, i.e., of origination of *Gonolkites* (from a species of *Parkinsonia*) and immigration of *Morphoceras*. The base of the Bathonian and the Zigzag Zone corresponds to the first occurrence level of *Gonolkites convergens* and the renewal of parkinsonids at the base of limestone bed RB071 (bed 23 in Sturani 1967) in the Ravin du Bès Section (Fig. 8). Additionally, the base of the Bathonian in Bas-Auran sections coincides with the lowest occurrence of *Morphoceras parvum*. Furthermore, the bases of the Northwest European primary standard Convergens Subzone and the Sub-Mediterranean secondary standard Parvum Subzone are in fact precisely coeval in the Bas-Auran area.

New palaeontological data about the youngest members of Bigotitinae and the oldest members of Zigzagiceratinae are of biochronostratigraphic importance for the subdivision and correlation of the basal Bathonian Zigzag Zone. Three successive biohorizons have been identified and chronocorrelated between the Bas-Auran (French Subalpine Basin) and Cabo Mondego (Lusitanian Basin) successions: Diniensis, Monde-

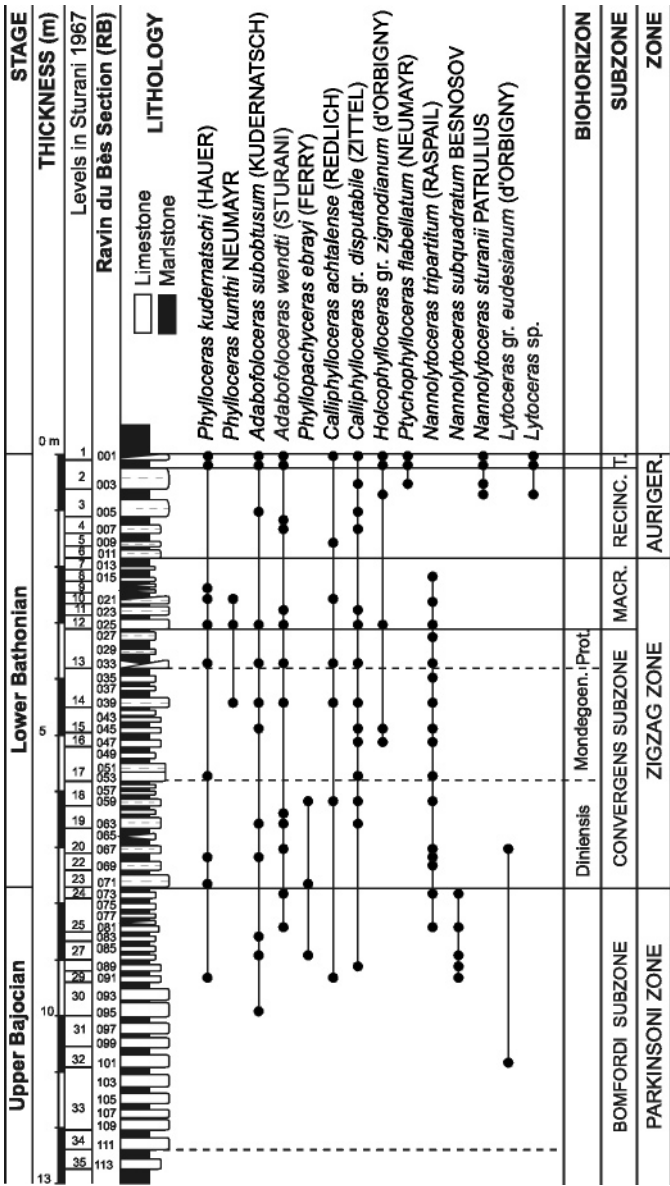


Fig. 7. Phylloceratina and Lytoceratina biochronostratigraphic data at the Bajocian/Bathonian boundary in the Ravin du Bès Section (from Pavia et al. 2008).

goensis and Protozigzagiceras biohorizons (Fernández-López et al. 2007a; Pavia et al. 2008). The Diniensis Biohorizon is characterized by the occurrence of *Bigotites diniensis* representatives and it corresponds to the lowest part of the Bathonian Zigzag Zone in the Sub-Mediterranean Province (e.g., Cabo Mondego and Bas-Auran). It encompasses the stratigraphic interval RB071–RB054 (Fig. 8, levels 23–18 of Sturani 1967) in Ravin du Bès Section. The Mondegoensis Biohorizon is defined by the lowest occurrence of *Bigotites mondegoensis* representatives. It comprises the stratigraphic interval RB053–RB034 (Fig. 8, levels 17–14 of Sturani 1967) in Ravin du Bès Section. The Protozigzagiceras Biohorizon is defined by the lowest oc-



According to Pavia et al. (2008), the record quality of the ammonoid biostratigraphic succession in the Bas-Auran area can be tested with various palaeontological criteria: the preservation state of fossil-specimens, taphonic populations and fossil-assemblages; abundance, concentration, packing and stratigraphic persistence of fossil-specimens; completeness, constancy and persistence of stratigraphic ranges; completeness and taxonomic diversity of successive fossil-assemblages; biostratigraphic turnover; proportion of virtual and actual palaeontological gaps in successive stratigraphic intervals; proportion of first and last occurrences of taxa; proportion of lowest and highest occurrences of taxa; successive or coincident clustering of last and first occurrences. Values of these twenty-one palaeontological attributes indicate a relatively homogeneous and good record quality, gradual biostratigraphic change and high degree of taxonomic similarity between the Bomfordi and Convergens subzones. These criteria, applied to the ammonoid genera which are known from the Bas-Auran area, also indi-

The Bajocian/Bathonian boundary may be characterized by secondary (auxiliary) biostratigraphic markers, such as nannofossils. According to the results of Erba (1990a, b; Cobianchi et al. 1992; Mattioli & Erba 1999), calcareous nannofossils are present in all beds and facilitate the characterization of the Bajocian-Bathonian transition. The Ravin du Bès Section appears to be suitable for the biostratigraphical study of microfossils.

such as foraminifers or ostracods, but there are at present no published studies. According to preliminary results (Bodergat in Mangold 1999), ostracods are present in all marly samples, but are badly preserved between bed RB071 and bed RB033. The marine taxa are different from those known in the Paris Basin and England. The Subalpine taxa, specially the genera *Pontocyprilla*, *Isobythocypris* and *Cordobairdia*, indicate deeper environments (more than 200 m). Palynomorphs are poorly preserved and are not yet stratigraphically useful across the boundary (Poulsen 1997; Mangold 1999).

Calcareous nannofossils (E. Erba & D. Tiraboschi)

Nannofossil biostratigraphic investigation was performed on 59 samples (approximately every 20 cm) collected from the Ravin du Bès section in the Bas-Auran area; sample figures correspond to the bed numbers of the lithostratigraphic column of Figures 7 and 8. This study is a revision of the previous work by Erba (1990a, b), extended to limestone layers and additional marlstone beds. Simple smear slides were prepared for both limestone and marlstone, using standard techniques without concentration by centrifuging cleaning in order to retain the original sediment composition. A few milligrams of powdered sediments were mounted on a glass slide with the Norland Optical Adhesive, and then analyzed using a polarizing light microscope, at 1250x magnification.

All studied samples contain calcareous nannofossils. A total of 37 taxa were identified and their distribution is given in Figure 9. The nannofossil total abundance fluctuates from extremely rare to common; the preservation is poor to moderate, with evidence of dissolution and overgrowth. Limestone levels generally contain depauperated and poorly preserved nannofloras, with stronger overgrowth and dissolution.

The nannofloras are characteristic of the Upper Bajocian–Lower Bathonian interval. Assemblages are dominated by *Watznaueria britannica* and *Watznaueria communis*, with common *Schizosphaerella punctulata*, *Watznaueria* aff. *W. manivitiae*, *Watznaueria manivitiae*, *Cyclagelosphaera margerelii*, *Cyclagelosphaera deflandrei*, *Lotharingius crucicentralis*, *Lotharingius velatus*, *Lotharingius sigillatus* and *Ethmorhabdus gallicus*.

Based on absence of *Carinolithus superbus* and of *Watznaueria barnesiae*, the lowermost portion of the investigated interval (samples 110 through 68b) corresponds to the Tethyan *W. communis* Subzone (NJT 10b) indicating a Late Bajocian age (Mattioli & Erba 1999). This subzone corresponds to the upper part of the Boreal NJ 10 Zone and the lower part of the NJ 11 Zone of Bown & Cooper (1998). The first occurrence (FO) of *Pseudoconus enigma* in sample 89 identifies the NJ10/NJ11 zonal boundary (Figs. 10–11). This taxon is rare and occurs only in limestones, with the only exception of a single specimen in marlstone sample 20, and this is why Erba (1990b) did not report this species.

The last occurrence (LO) of *Hexalithus magharensis* was observed in sample 82 indicating a latest Bajocian age (Mattioli

& Erba 1999). Similarly, Erba (1990b) recorded this event in the Parkinsoni Zone (latest Bajocian) of the Digne area, whereas in Portugal and Morocco de Kaenel et al. (1996) found an older age for the LO of *H. magharensis*, calibrated between the end of the Early Bajocian and the beginning of the Late Bajocian.

The FO of *Stephanolithion speciosum octum* was observed in sample 76; the taxon is extremely rare and scarce in the studied section. This event has been correlated to the base of the Parkinsoni Zone in NW Europe and Portugal (de Kaenel et al. 1996), but within the Zigzag Zone in SE France (Erba 1990b). Bown et al. (1988) and Bown & Cooper (1998) report the FO of *S. speciosum octum* at the base of the Boreal NJ 11 Zone.

The FO of *W. barnesiae* (NJT11) was observed in sample 68a of earliest Bathonian age (Zigzag Zone). This event defines the base of the Tethyan NJT11 Zone (Mattioli & Erba 1999), comparable to most of the Boreal NJ11 Zone and NJ12a Subzone (Bown et al. 1988, Bown & Cooper 1998).

The uppermost portion of the studied interval corresponds to the Tethyan NJT 11 Zone (Mattioli & Erba 1999), since *Cyclagelosphaera wiedmannii* was not observed.

From sample 89 upwards, rosette-shaped specimens similar to the genus *Rucinolithus* were consistently observed. They show highest abundance in the interval between sample 45 through 22 (Fig. 9), both in limestone and marlstone beds. Two morphotypes were distinguished, namely small (<8 microns) and large (>8 microns) *Rucinolithus* spp., based on their diameter. More detailed investigations are in progress to characterize the taxonomy of these morphotypes (Tiraboschi & Erba in prep.).

Our results are consistent with previous biostratigraphic data from the Upper Bajocian–Lower Bathonian interval in SE France (Erba 1990b), Portugal, NW Europe (de Kaenel & Bergen 1993, de Kaenel et al. 1996), Lombardian Basin (Chiari et al. 2007) and Boreal Realm (Bown & Cooper 1998). For the first time *P. enigma* has been documented from mid to low latitudes allowing a direct calibration between Tethyan and Boreal nannofossil events and biozones (Fig. 10).

Correlation (S.R. Fernández-López)

Ammonites are the most relevant taxonomic group for global biochronostratigraphic correlation of the Bajocian/Bathonian boundary. Nevertheless, diverse taxonomic groups of macroinvertebrates and microfossils are also of biochronostratigraphic relevance.

Ammonites

Late Bajocian and Early Bathonian ammonites are found worldwide in the three major, oceanic or marine, palaeogeographical units: Tethyan, Pacific and Boreal domains or realms (Cariou et al. 1985; Hillebrandt et al. 1992a, b; Taylor et al. 1992; Westermann 1993a, 2000; Page 1996a; Enay & Cariou 1999; Moyne & Neige 2007). The most difficult problem in biochronocorrelation of the boundary is not the low biostratigraphic

Lithology	Samples	Total abundance	Preservation	Species
Limestone	1	FC	PM	<i>Watznaurea communis</i> Reinhardt
Limestone	3a	R	P	<i>Watznaurea britannica</i> (Stradner) Reinhardt
Limestone	3b	R	P	<i>Watznaurea fossacincta</i> (Black) Bown in Bown & Cooper
Marlstone	4	FC	PM	<i>Watznaurea contracta</i> (Bown & Cooper) Cobianchi et al.
Limestone	5	R	P	<i>Watznaurea aff. W. manivillae</i> Cobianchi et al.
Marlstone	6	FC	PM	<i>Hexalithus magharensis</i> Mostkowitz & Ehrlich
Limestone	7	R	P	<i>Lotharingius hauffii</i> Grün and Zwelli in Grün et al.
Marlstone	8	FC	PM	<i>Cyclagelosphaera magerelii</i> Noël
Limestone	9	R	P	<i>Cyclagelosphaera deflandrei</i> (Manivth) Roth
Marlstone	10	FC	PM	<i>Dicormatolithus lehrmanni</i> Noël
Marlstone	12	F	PM	<i>Zeugmatolithus erectus</i> (Deflandre in Deflandre & Fert) Reinhardt
Limestone	13	R	P	<i>Dicormatolithus</i> sp.
Marlstone	14	R	P	<i>Schizosphaerella punctulata</i> Deflandre & Dangeard
Marlstone	16	R	P	<i>Triscutum</i> sp.
Marlstone	20	FC	PM	<i>Tubirhabdus patulus</i> Pnns ex Rood et al.
Limestone	21	R	P	<i>Watznaurea aff. W. contracta</i> Cobianchi et al.
Marlstone	22	FC	PM	<i>Lotharingius sigillatus</i> (Stradner 1961) Pnns in Grün et al.
Limestone	23	R	P	<i>Lotharingius velatus</i> Bown and Cooper
Marlstone	24	FC	PM	<i>Ethmorhabdus gallicus</i> Noël
Limestone	25	R	P	<i>Lotharingius cruciellensis</i> (Medd) Grün and Zwelli
Marlstone	26	FC	PM	<i>Tetralithus</i> sp.
Limestone	31	R	P	<i>Lotharingius barozii</i> Noël
Marlstone	34	F	PM	<i>Watznaurea aff. W. barnesiae</i> – with central area not completely closed
Limestone	35	R	P	<i>Hexalithus hexalithus</i> Noël
Marlstone	39	RF	P	small <i>Rucinolithus</i> sp.
Marlstone	40	F	P	large <i>Rucinolithus</i> sp.
Limestone	45	RF	P	<i>Pseudocornus enigma</i> Bown & Cooper
Marlstone	46	F	P	<i>Zeugmatolithus</i> sp.
Limestone	47	RF	P	<i>Stephanolithon speciosum octum</i> Rood & Barnard
Marlstone	48	FC	PM	<i>Hexapodirhabdus curvillieri</i> Noël
Limestone	51	RF	P	<i>Watznaurea barnesiae</i> (Black in Black & Barnes) Perch-Nielsen
Limestone	53	RF	P	<i>Dicormatolithus striatus</i> Mostkowitz & Ehrlich
Limestone	55	R	P	<i>Biscutum</i> sp.
Marlstone	60	C	M	<i>Truncatocapulus hexaporus</i> Mostkowitz & Ehrlich
Limestone	63	RF	P	<i>Crepidolithus crassus</i> (Deflandre in Deflandre & Fert) Noël
Marlstone	64	FC	PM	<i>Stephanolithon speciosum speciosum</i> Deflandre in Deflandre & Fert
Limestone	67	R	P	
Marlstone	68a	C	PM	
Marlstone	68b	C	PM	
Limestone	69	R	P	
Limestone	71a	R	P	
Limestone	71	R	P	
Limestone	71b	F	P	
Marlstone	72	C	M	
Limestone	73	R	P	
Marlstone	74	C	PM	
Marlstone	76	C	M	
Marlstone	82	C	PM	
Limestone	83	R	P	
Marlstone	84	C	M	
Limestone	85	F	PM	
Marlstone	86	FC	P	
Limestone	89	F	P	
Marlstone	90	C	PM	
Marlstone	92	C	PM	
Limestone	93	FC	P	
Marlstone	96	C	PM	
Marlstone	100	C	PM	
Marlstone	102	C	PM	
Marlstone	110	C	PM	

Fig. 9. Chart of calcareous nannofossil ranges of the Ravin du Bès Section. The semiquantitative distribution of nannofloras was coded as follows: Total abundance. C = 10–15 specimens in each field of view; FC = few/common: 5–9 specimens in each field of view; F = few: about 5 specimens in each field of view; RF=rare/few: 2–5 specimens in each field of view; R = rare: 1 specimen in each field of view. Preservation. M = moderate: little evidence of dissolution and/or overgrowth is present; primary morphological characteristics are sometimes altered; MP = moderate/poor: evidence of dissolution and/or overgrowth is present; primary morphological characteristics are sometimes modified, fragmentation has occurred; P = poor: most specimens exhibit dissolution or overgrowth; primary morphological characteristics are sometimes destroyed; fragmentation has occurred. Species abundance. A = abundant: >10 specimens in each field of view; C = common: 1 specimen in 1–9 fields of view; FC = few/common: 1 specimen every 10 fields of view; F = few: 1 specimen in 11–29 fields of view; RF = rare/few: 1 specimen every 30 fields of view; R = rare: 1 specimen in 31–100 fields of view.

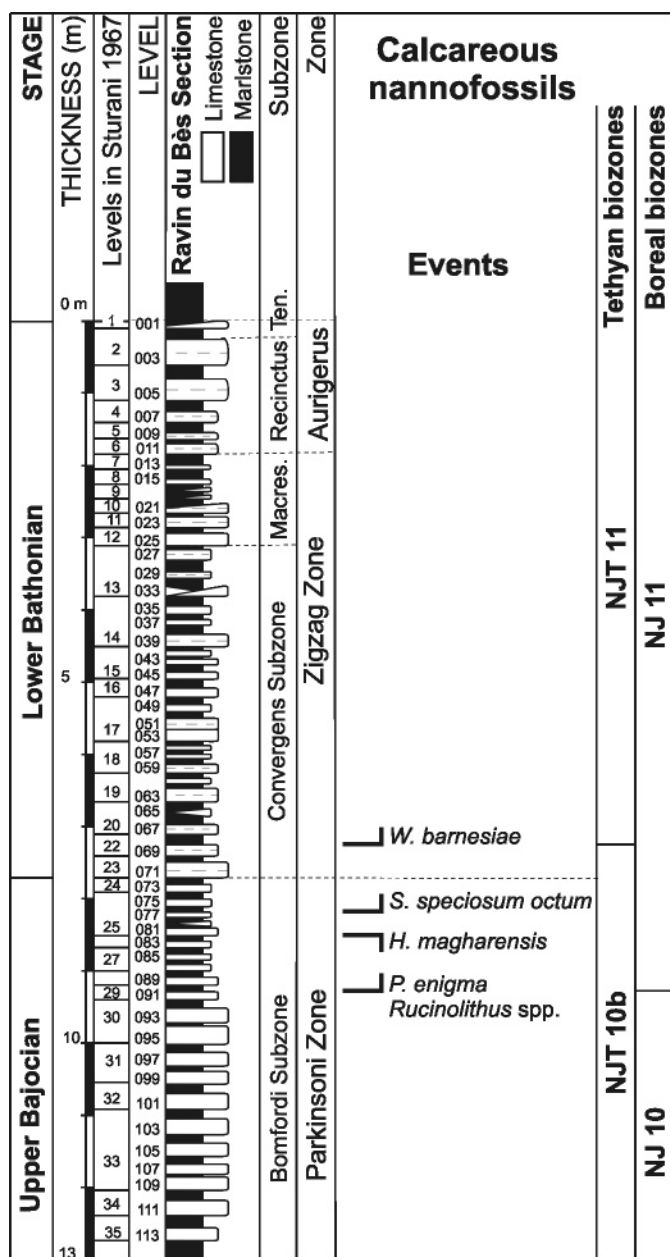


Fig. 10. Nannofossil events detected in the Ravin du Bès Section. Tethyan biozones after Mattioli & Erba (1999) and Boreal biozones after Bown & Cooper (1998).

turnover of the ammonoid succession across the boundary in Bas Auran area or the low faunal turnover at the Bajocian/Bathonian transition, but the strong provincialism with three separate realms.

Figure 12 shows standard zonations for the three ammonite biogeographical provinces represented in western Europe. Ammonites of the Zigzag Zone have a wide distribution through the Northwest European, Sub-Mediterranean and Mediterranean provinces of the West Tethyan Subrealm. In the Bas-Auran area, where Northwest European and Sub-Mediterranean taxa are relatively common, the primary standard Convergents Subzone and the secondary standard Parvum Subzone can be recognized.

The Northwest European Province, in which parkinsonids are common, comprises the following epeiric areas: England (Torrens 1980; Callomon 1995, 2003; Callomon & Cope 1995; Page 1996b, 2001; Dietze and Chandler 1998; Chandler et al. 1999), Normandy, Boulonnais, Lorraine, Alsace, northern Jura (Mangold & Rioult 1997, Rioult et al. 1997, Thierry 2003), northern Germany (Westermann 1958; Metz 1990, 1992), northern and central Poland (Kopik 2006; Zaton 2006).

The Sub-Mediterranean Province, in which Bathonian *Morphoceras* [M] – *Ebrayiceras* [m] occur associated with parkinsoniids and scarce phylloceratids and lytoceratids, includes the following epeiric areas: Lusitanian Basin (Fernández-López et al. 2006a, b), Iberian Basin (Fernández-López 2000, 2001), Aquitaine, Causses, Centre-west France, Nièvre (Delance et al. 1979; Courville et al. 1999; Enay et al. 2001), Mâconnais, Ardèche, southern Jura (Elmi 1967; Mangold 1971a, b, c, 1997a, b; Rulleau 2006), western Alps and Subalpine Basin (Sturani 1967; Pavia & Sturani 1968; Pavia 1973, 1984; Torrens 1987; Innocenti et al. 1990; Zany et al. 1990; Joly 2000), southern Germany (Dietl 1978, 1981, 1982, 1983, 1986, 1988; Dietl et al. 1978, 1983; Dietl & Hugger 1979; Dietl & Kapitzke 1983; Callomon et al. 1987; Schairer 1987, 1994; Dietze & Chandler 1996; Köstler & Schairer 1996; Dietze et al. 1997, 2002, 2004, 2007; Schweigert & Dietze 1998; Dietze 2000; Dietze & Schweigert 2000; Schweigert et al. 2002, 2003, 2007; Ohmert et al. 2004; Dietze & Dietl 2006), South Poland (Luczynski et al. 2000; Matyja & Wierzbowski 2000, 2001; Zaton & Marynowski 2006), Pieniny Klippen Belt (Wierzbowski et al. 1999; Schlögl & Rakús 2004; Schlögl et al. 2005, 2006), South Transdanubian Mecksek (Galács 1995a; Geczy & Galács 1998), Romania (Galács 1994; Patrulius 1996), Balkans (Stephanov 1972), northwestern-central Iran (Seyed-Emami et al. 1985, 1989, 1991, 1998a, b) and north-eastern Iran (Majidifard 2003).

The Mediterranean Province, in which Late Bajocian and Early Bathonian morphoceratids occur associated with common phylloceratids and lytoceratids, comprises the following shelfal or oceanic areas: Betic Basin (Mangold 1981; Sandoval 1983, 1986, 1990; Sandoval et al. 2001), Majorca (Sandoval 1994), Sicily (Wendt 1963, 1971; Galács 1985, 1999a, b; Pavia & Cresta 2002; Pavia et al. 2002; Martire & Pavia 2004; Pavia 2007), Alps (Sturani 1971; Krystyn 1972; Joly 2000; Martire 1989; Mangold & Gygi 1997) and North Transdanubian Hungary (Galács 1980, 1993, 1995b).

In the north-eastern Tethyan border (Donetz, Crimea, Caucasus, Great Balkan, Turkmenistan, Tadjikistan, Uzbekistan, Kazakhstan) latest Bajocian to Early Bathonian parkinsonids and morphoceratids have been described, below Middle Bathonian specimens of *Bullatimorphites*, *Procerites* and *Siemiradzkia* (Rostovtsev 1985; Tseretely 1989; Beznosov & Mitta 1998, 2000).

In Tibet and South-East Asia, Bathonian *Siemiradzkia*, *Procerites* and *Wagnericeras* have been identified (Cariou & Enay

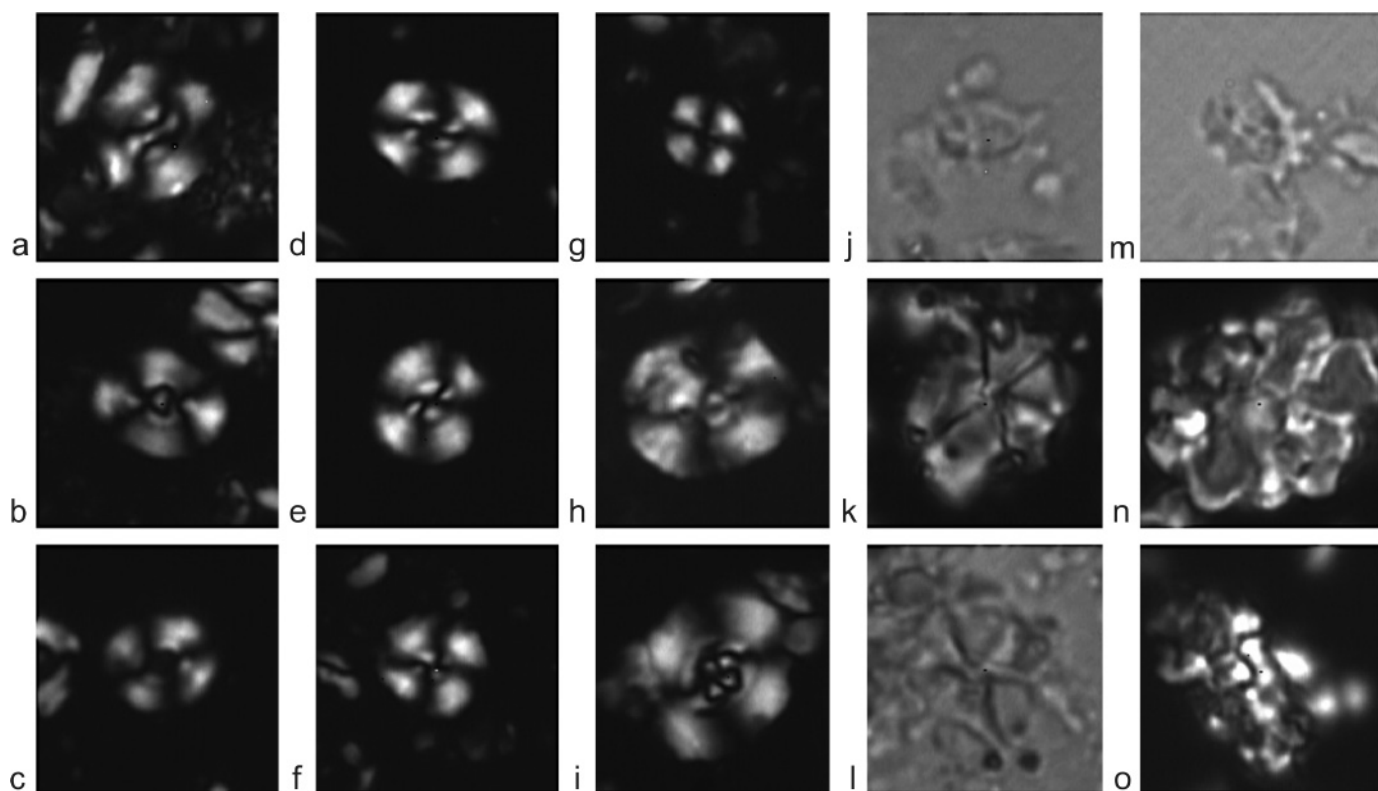


Fig. 11. Upper Bajocian and Lower Bathonian calcareous nannofossils from the Bas-Auran area. All specimens at 1250x magnification. (a): *Watznaueria communis*, crossed nicols, sample 110. (b): *Watznaueria britannica*, crossed nicols, sample 102. (c): *Watznaueria fossacincta*, crossed nicols, sample 110. (d): *Watznaueria* aff. *W. barnesiae*, crossed nicols, sample 60. (e): *Watznaueria barnesiae*, crossed nicols, sample 68a. (f): *Watznaueria barnesiae*, crossed nicols, sample 10. (g): *Cyclagelosphaera margerelii*, crossed nicols, sample 40. (h): *Watznaueria manivitiae*, crossed nicols, sample 67. (i): *Watznaueria* aff. *W. contracta*, crossed nicols, sample 48. (j): *Stephanolithion speciosum octum*, transmitted light, sample 76. (k): *Hexalithus magharensis*, crossed nicols, sample 110. (l): *Hexalithus magharensis*, transmitted light, sample 110. (m): *Stephanolithion speciosum speciosum*, transmitted light, sample 76. (n): *Rucinolithus* sp., crossed nicols, sample 6. (o): *Pseudoconus enigma*, crossed nicols, sample 53. Scale bar is 5 μ m.

NW European Province			Sub-Mediterranean Province		Mediterranean Province		
England, Normandy, Boulonnais, Lorraine, Alsace, northern Germany, northern and central Poland.			Portugal, Iberian Basin, Aquitaine, Causses, Centre-west France, Nièvre, southern Jura, Mâconnais, Ardèche, southern Germany, southern Poland, Balkans, northern and central Iran.		Betic Basin, Sicily, Appennines, Switzerland, Austria, Hungary p.p. (Villany and Mecsek), Hellenids, Serbia.		
Lower Bathonian	Tenuiplicatus		Aurigenus	Tenuiplicatus		Aurigenus	Postpollubrum
	Zigzag	Yeovilensis		Recinctus			Yeovilensis
		Macrescens	Zigzag	Macrescens	Zigzag	Macrescens	
		Convergens		Parvum		Dimorphitiformis	
Upper Bajocian	Parkinsoni	Bornfordi	Parkinsoni	Bornfordi	Parkinsoni	Dimorphus	
		Truellei		Densicosta		Daubenyi	
		Acris		Acris			

Fig. 12. Ammonite zones and subzones of the Uppermost Bajocian and Lower Bathonian in several palaeobiogeographical provinces: Northwest European (Westermann & Callomon 1988; Callomon & Cope 1995; Callomon 2003), Sub-Mediterranean (Mangold 1990; Rioult et al. 1997; Mangold & Rioult 1997) and Mediterranean (Galáz 1980, 1993; Sandoval 1983, 1990; Sandoval et al. 2001; O'Dogherty et al. 2006) provinces.

1999; Yin et al. 2000; Yin 2005). Upper Bajocian *Leptosphinctes* and *Cadomites* have been recognized in Japan, below Upper Bathonian *Pseudoneuquenicer* (Sato 1992).

Lower Bathonian morphoceratids, parkinsoniids and phylloceratids have been recognized in several basins of the southern Tethyan border: Morocco, Algeria and Tunisia (Elmi 1971; Elmi & Alméras 1984; Enay et al. 1987b; Ouahabi 1994; Soussi et al. 2000). *Oranicer*, *Micromphalites* and *Oxycer* occur in Lower Bathonian, whereas *Ermoceras*, *Leptosphinctes* and *Oppelia* characterize Upper Bajocian deposits.

In the Ethiopian Province, including Libya, Egypt, Israel, southern Turkey, southern Iran and Saudi Arabia (Parnes 1981, 1984, 1988; Enay et al. 1986, 1987a; Enay & Mangold 1994, 1996), the Lower Bathonian comprises the Tuwaiqensis (*Tu-lites*) and Clydocromphalus (*Micromphalites*) zones of the Arabian Province, including *Procer* and *Zigzag* in basal levels. Upper Bajocian deposits with *Ermoceras*, *Thambites*, *Leptosphinctes* and *Spirocer* represent the Runcinatum (*Ermoceras*), Mogharens (*Ermoceras*) and Planus (*Thambites*) zones.

In the Indo-Malgach Province, Late Bajocian and Middle Bathonian ammonites, but not Early Bathonian, have been described from Kenya, Madagascar and India (Singh et al. 1982, 1983; Jaitly & Singh 1983, 1984; Pandey & Agrawal 1984; Pandey & Westermann 1988; Gal  cz 1990; Pandey & Callomon 1995; Prasad et al. 2007; Roy et al. 2007).

In south-western Pacific areas, Sula, Irian Jaya and New Guinea, latest Bajocian to Early Bathonian *Pretulites* and Early to Middle Bathonian *Satoceras*, as well as Bathonian specimens of *Asphinctites*, *Rugiferites* and *Bullatimorphites* have been described (Westermann & Getty 1970; Westermann & Callomon 1988; Sukanto & Westermann 1992; Westermann 1995, Callomon & Rose 2000).

Separate Late Bajocian and Early Bathonian ammonoid faunas have been distinguished, associated with characteristic Eurycephalitinae, in the southern East-Pacific Subrealm of the Tethyan Realm: New Zealand (Westermann & Hudson 1991; Westermann 1993b; Westermann et al. 2000, 2002), Argentina, Chile and Peru (Westermann & Riccardi 1980; Westermann et al. 1980; Riccardi 1985, 1991; Riccardi et al. 1990a, b, 1991, 1992, 1994; Riccardi & Westermann 1991a, b, 1999; Hillebrandt et al. 1992a, b; Fern  ndez-L  pez et al. 1994; Gr  schke & Hillebrandt 1994; Hillebrandt 1995, 2001; Gr  schke 1996; Parent 1998). *Leptosphinctes*, *Lupherites*, *Strenoceras*, *Spiroceras* and *Megasphaeroceras* occur in the Upper Bajocian of the Andean Province. *Lobosphinctes intersertus* Buckman has been identified in Chacay Melehue (Argentina) below a Bathonian *Cadomites*-Tulitidae mixed assemblage. The first occurrence of several genera such as *Oxycerites*, *Zeissoceras*, *Prohecticoceras* and *Rugiferites*, below the oldest representatives of Bathonian *Bullatimorphites*, have been used to recognize Lower Bathonian deposits. New species of Bathonian ?*Zigzagiceras* and *Morphoceras* have been proposed (Gr  schke & Hillebrandt 1994; Riccardi & Westermann 1999).

In Mexico (Sandoval & Westermann 1986; Sandoval et al. 1990) Upper Bajocian begins with the upper Flores Zone of Oaxaca, containing the Mediterranean *Subcollina lucretia* (Orbigny). The overlaying Zapotecum Zone includes *Parastrenoceras*, *Leptosphinctes* and *Oppelia*. The Upper Bathonian Retrocostatum Zone has been identified by *Prohecticoceras blanazense*, associated with *Epistrenoceras*, *Lilloettia* and *Neuquenicerias*.

In the Western Interior of the United States of America (Shoshonean Province, Imlay 1981), western Canada and southern Alaska (Athabaskan Province; Imlay 1980, 1982, 1984; Hall & Westermann 1980; Hall & Stronach 1981; Callomon 1984; Hall 1984, 1988, 1989; Poulton et al. 1991, 1994) the Upper Bajocian Rotundum Zone includes *Leptosphinctes*, *Lupherites*, *Spiroceras* and *Megasphaeroceras*, below the *Epizigzagiceras-Parareineckia* association. The *Parachondrocera-Sohlites* assemblages from Oregon may be Upper Bajocian or Lower Bathonian (Imlay 1984).

The Boreal Realm (Eastern Greenland, Siberia, Northern Alaska and Northern Canada) became clearly differentiated in the Late Bajocian and several zonations for the Early Ba-

thonian have been proposed. The Cardioceratidae, in particular *Cranocephalites* and *Arctocephalites*, constituted characteristic elements of the Boreal Realm at the Bajocian/Bathonian boundary (Callomon 1985). The Zone of *Arctocephalites arcticus* (NEWTON & TEALL), above the Zone of *Cranocephalites pompeckji* (MADSEN), may represent the basal Bathonian zone in the Boreal Realm (Callomon 1993, 1994, 2003; Rawson 1982; Zakharov et al. 1998). The Zone of *Arctocephalites spathi* from northern Yukon probably is coeval with the Boreal Articus Zone of eastern Greenland (Poulton 1987). Boreal Arctocephalitinae are associated with parkinsoniids in the south-eastern part of the Russian platform, allowing the correlation between the regional Michalskii-Besnosovi zonal boundary and the Boreal Arcticus-Greenlandicus boundary or the Northwest European Parkinsoni-Zigzag boundary (Mitta 2001, 2004, 2005, 2006, 2007; Mitta & Seltzer 2002; Mitta et al. 2004; Saltykov 2007; Zakharov 2007).

Other taxonomic groups

Several authors have proposed diverse biozonations for the Upper Bajocian and Lower Bathonian based in different taxonomic groups of macroinvertebrates: brachiopods (Mance  ido & Dagys 1992; V  r  s 2001; Alm  ras et al. 2007), belemnites (Challinor 1992; Challinor et al. 1992; Comb  morel 1997), nautiloids (Branger 2004), bivalves (Damborenea et al. 1992; Hallam 1994; Damborenea 2002; Ruban 2006), echinoderms (Thierry et al. 1997; Moyne et al. 2005), corals (Beauvais 1992).

The following taxonomic groups of microfossils are of biochronostratigraphic relevance also: foraminifera (Bassoulet 1997; Ruget & Nicollin 1997; Gr  fe 2005; Cai et al. 2006; Saltykov 2007; Wernli & G  r  g 2007), ostracods (Braun & Brooke 1992; Bodergerat 1997), dinoflagellate cysts (Riding & Thomas 1992; Fauconnier 1997; Poulsen & Riding 2003), radiolarians and calcareous nannofossils (Pessagno & Mizutani, 1992; Baumgartner et al. 1995; Cordey et al. 2005; Chiari et al. 2007). Palaeobotanical and palynological data have been recently published by: Kimura et al. 1992; Sarjeant et al. 1992; Cleal & Rees 2003; Wang et al. 2005; Vaez-Javadi & Mirzaei-Ataabadi 2006; Jana & Hilton 2007.

Isotope stratigraphy

From a geochemical point of view, in the French Subalpine Basin during the Jurassic Period, several authors have emphasized that the manganese content of pelagic carbonates is related to 2nd-order sea-level changes and episodes of hydrothermal activity that affected the chemistry of global sea water. The main transgressive phases are marked by a manganese content increase, whereas regressive phases are characterized by decreasing trends (Corbin 1994; Corbin et al. 2000). In the Chaudon-Norante section, 4 km north of the Bas-Auran area, the Early Bathonian maximum transgressive is marked by sedimentary condensations, associated with high manganese content (from 300 to 1370 mg kg⁻¹). In contrast, the Middle and

Late Bathonian regressive phase coincides with low manganese content periods. These stratigraphical patterns in divalent manganese can be of either local or regional significance, being concentrated, most probably as a very early diagenetic phase, only in oxygen-depleted waters that typically underlie zones of elevated organic productivity (Jenkyns et al. 2002). For strontium isotope ($^{87}\text{Sr}/^{86}\text{Sr}$ ratio), oxygen isotope ($\delta^{18}\text{O}$) or carbon isotope ($\delta^{13}\text{C}$) chemostratigraphy, no data are currently available.

Volcanogenic deposits suitable for direct radio-isotope dating are not known in the section. The age of the Bajocian/Bathonian boundary has been dated 167.7 ± 3.5 Ma by Gradstein & Ogg (2004) and Gradstein et al. (2005).

Magnetostratigraphy (R. Lanza)

In the autumn 1994, the Bas-Auran section was extensively sampled across the Bajocian/Bathonian boundary. Some 30 hand-samples were collected and cored in the rock-magnetism laboratory of Torino University. All specimens were measured with a JR-5 spinner magnetometer and thermally demagnetized using a Schonstedt furnace (Degiorgis 1996).

The specimens are characterized by three remanent magnetization components: (1) a viscous (VRM) component close to the present field and removed at temperature values about 120 °C; (2) a secondary component stable up to values of 300 to 450 °C; and (3) a high-temperature component, stable between 350–450 and 480–500 °C.

The secondary component has been interpreted as a Tertiary magnetic overprint: declination points to the NW (310° to 330°), inclination is positive (30° to 50°).

The high-temperature component has been isolated only in 11 out of the 32 analyzed specimens. Its definition is difficult and often poor, because its intensity is very low, usually

10–25% of the initial NRM. This component may be regarded as the more stable fraction of the primary Jurassic remanence acquired when the rocks formed. The fact that the primary remanence can be isolated only in few specimens and its poor definition prevent any reliable magnetostratigraphic interpretation.

The results found by Degiorgis (1996) in the Bas-Auran section have been fully substantiated in the whole southern Subalpine Chains by Aubourg & Chabert-Pelline (1999).

Gamma-ray spectrometry

Field gamma-ray spectrometry data have been obtained by G. Pavia, P. Lazzarin and L. Leroy (April 2007) and are shown in Figure 13. Spectral gamma-ray data from the Ravin du Bès Section show an increase in the total gamma-ray counts at the Aurigerus Zone. The values are relatively low and display insignificant variation at the Bajocian-Bathonian boundary, but they show a positive peak at the top of the Lower Bathonian. Total gamma-ray logs have been used in sequence stratigraphy on the basis that gamma-ray peaks commonly correspond to maximum flooding surfaces (cf. Parkinson 1996; Deconinck et al. 2003; Pawellek & Aigner 2003, 2004; Pellenard et al. 2003; Raddadi et al. 2005; Ruf et al. 2005; Schnyder et al. 2006). High gamma-ray counts, low sedimentation rates and high concentrations of ammonites may be associated with the development of condensed sections in carbonate environments. These features, however, developed both in condensed deposits of deep carbonate environments during transgressions or episodes of relative sea-level rise and in expanded deposits of shallow carbonate epicontinental platforms during regressions or episodes of relative sea-level fall (Fernández-López et al. 2002). The stratigraphic trend in spectral gamma-ray data associated with sedimentary condensation on the Bas-Auran area, from the

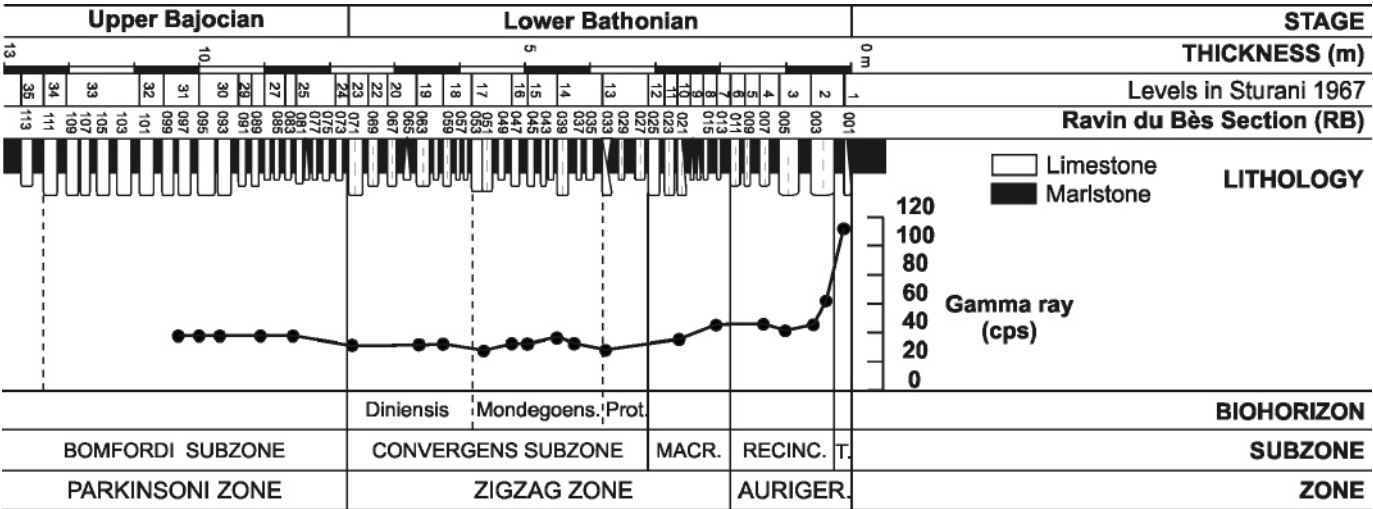


Fig. 13. Gamma-ray log of the Ravin du Bès Section showing a positive peak at the top of the “Marno-calcaires à *Cancellophycus*” in the Lower Bathonian Tenuiplicatus Subzone (Aurigerus Zone).

Bajocian Bomfordi Subzone towards Bathonian *Tenuiplicatus* Subzone, provides support for an Early Bathonian deepening half-cycle of 2nd order, lacking evidence of stratigraphic gaps at the Bajocian-Bathonian transition.

Therefore, the current data do not support the existence of a gap of the order of a whole biohorizon or even a subzone. The base of the bed RB071 is a minor sedimentary and stratigraphic discontinuity (tested with sequence stratigraphy and sedimentological data, as well as with palaeoichnological and taphonomic analyses), lacking evidence of significant hiatus (such as a peak in the spectral gamma-ray data shown in Fig. 13), non-existing evidence of biostratigraphic gap (as argued with diverse criteria in the chapter of record quality by Pavia et al. 2008) or missing biochronostratigraphic unit (the first subzone at the base of the Zigzag Zone, with three successive biohorizons, shows the highest biochronostratigraphic completeness, so far only recognized in the Cabo Mondego and Bas Auran sections, Fernández-López et al. 2007a).

Site protection (M. Guiomar)

The sites of Bas-Auran and Le Bès are part of the protected territory of “La Réserve Naturelle Géologique de Haute Provence”. The Geological Reserve covering 55 communes in the departments “Alpes de Haute-Provence” and “Var” has been entrusted with the main missions of protecting, enhancing and raising awareness of the environment and supporting economic development of this heritage. These different missions are based on scientific knowledge of geological sites and inventories. This is why the Reserve wishes to encourage the number of scientific studies within its territory. While some of this research is instigated and supported by the Reserve itself and conducted in collaboration with its partners, academic or others, there is also independent research, conducted with technical support from the Reserve, in particular for the authorization to collect fossils. Two types of regulations apply in the protected territory: those relative to sites listed as natural reserves (Réserve Naturelle) by ministerial decree and those relative to protected areas, by prefectural order (arrêtés préfectoraux). This explains why the sites of Bas-Auran and Le Bès do not have the same status. Bas-Auran is listed as a natural reserve (RN): no collecting or surface removal is possible, except by ministerial authorization. The section of Le Bès is part of the protected area where fossils may be collected and authorizations delivered for extraction or excavation (files managed by the Reserve), mainly in the context of research projects. Authorizations delivered by the Reserve in no way affect private property rights and all applications for excavation must be accompanied by a request to the owner of the land. According to the site’s scientific value, the Reserve may request that a geological site be added to or removed from listing to facilitate its management and/or protection; exceptionally, the Reserve may approach local communities for the acquisition of certain lands. For the section(s) to obtain GSSP status, it is necessary to undertake all the requisite measures to maintain free access

to them as well as consider minimum developments (ease of access, safety) as well as enhancement (to be defined with scientists and managers), whenever possible.

The Bajocian/Bathonian auxiliary stratotype in Cabo Mondego Section, Portugal (S.R. Fernández-López, M.H. Henriques & C. Mangold)

An auxiliary section and point (hereafter ASSP) for the base of the Bathonian Stage is located in Cabo Mondego, 40 km west of Coimbra, 7 km north of Figueira da Foz (40° 11' 19" N, 8° 54' 30" W, Section 02). It provides complementary data about the ammonite succession and biochronostratigraphic subdivision of the Sub-Mediterranean Parvum Subzone and the Northwest European Convergens Subzone.

These classical fossiliferous deposits have been studied by numerous specialists (Ruget-Perrot 1961; Elmi et al. 1971; Mangold 1971c, 1990; Rocha et al. 1981, 1987; Mangold & Rioult 1997; Fernández-López & Henriques 2002; Fernández-López et al. 2006a, b). The Bathonian deposits correspond to the Cabo Mondego Formation and comprise limestone-marl alternations, with ammonoids, bivalves (*Bositra*), rhynchonellid brachiopods, crinoids and belemnites. Bioturbation is common (*Zoophycos*, *Thalassinoides*, *Chondrites*). These fossiliferous deposits were developed in an open sea, in hemipelagic environments of distal and outer carbonate ramp, below the fair weather wave base (Watkinson 1989; Soares et al. 1993; Azerêdo et al. 2003).

In the Lusitanian Basin, Upper Bajocian and Lower Bathonian Phylloceratina and Lytoceratina account for less than 1% of the total ammonoid assemblage, and parkinsonids are very scarce (less than 5.0%). Successive ammonoid fossil assemblages are composed of Submediterranean taxa, but they allow correlation with the zonal scales of the diverse basins of the Mediterranean and NW European provinces.

The base of the Bathonian has been established by the lowest occurrence of representatives of the *Morphoceras* [M] – *Ebrayiceras* [m] group, at the base of the marly interval 123 of Section-02 (Fig. 14), which corresponds to the base of the marly interval FC1 of Section-90. The Lower Bathonian index ammonite *Morphoceras parvum* occurs in the marly interval 02CM139. The Lower Bathonian index ammonite *Gonolkites convergens* occurs in the marly interval 02CM181. From a biochronostratigraphic point of view, 10 metres of thickness with 62 successive ammonoid fossil-assemblages from 77 successive fossiliferous stratigraphic intervals have been recognized and sampled in the Parvum Subzone.

New taxa of Perisphinctidae, based on data from Cabo Mondego, are of primary relevance for the biochronostratigraphic subdivision and correlation of the Submediterranean Parvum Subzone. The lowest occurrences of *Bigotites mondegoensis* and *Protozigzagiceras* correspond to two successive biostratigraphic events allowing distinction of three successive biohorizons in the Parvum Subzone in Cabo Mondego (Lusitanian Basin) and Bas Auran (Alpine Basin): the Di-niensis, Mondegoensis and Protozigzagiceras biohorizons

du Bès Section belonging to three biohorizons of the Parvum Subzone have been recognized. The Bomfordi Subzone attains a minimum thickness of 5 m and includes 42 successive ammonoid fossil-assemblages.

The boundary has been characterized by both primary and secondary (auxiliary) biostratigraphic markers. There is a well-preserved, abundant and diverse fossil record across the boundary interval, with key markers (ammonites and nannofossils) for worldwide correlation of the uppermost Bajocian and Lower Bathonian. The section appears to be suitable for biostratigraphic study of microfossils, such as foraminifera, but as yet there are no published studies.

Regional analyses of sequence stratigraphy and manganese chemostratigraphy are available. A transgressive systems tract associated with a deepening phase and sedimentary starvation, within 3rd and 2nd order deepening/shallowing cycles, was developed in the Bas-Auran area of the French Subalpine Basin, during the Early Bathonian. No data are currently available for strontium isotope ($^{87}\text{Sr}/^{86}\text{Sr}$ ratio), oxygen isotope ($\delta^{18}\text{O}$) or carbon isotope ($\delta^{13}\text{C}$) chemostratigraphy.

The stratigraphic trend in spectral gamma-ray data provides support for an Early Bathonian deepening half-cycle of second order.

Bajocian and Bathonian deposits have been remagnetized with a steady normal polarity. The requirement of suitability for magnetostratigraphy and geochronometry, however, can be indirectly satisfied by reference to the Bathonian magnetostratigraphic scale of Steiner et al. (1987; O'Dogherty et al. 2006) as defined in the Subbetic Cordillera.

Volcanogenic deposits suitable for direct radio-isotope dating are not known in the section. According to the data published by Gradstein & Ogg (2004) and Ogg (2004), the age of the Bajocian/Bathonian boundary is 167.7 ± 3.5 Ma in other basins.

The criteria of accessibility, conservation and protection are assured by the "Réserve Naturelle Géologique de Haute Provence", protected under national law and recognised by UNESCO. The park is managed by the "Centre de Géologie de Digne".

The Cabo Mondego Section is suggested as the Bathonian ASSP within the same GSSP proposal. It provides complementary data about the ammonite succession and biochronostratigraphic subdivision of the Sub-Mediterranean Parvum Subzone and the Northwest European Convergens Subzone, at the basal Bathonian Zigzag Zone. Accessibility, conservation and protection are guaranteed, after the classification of the Cabo Mondego area as a Natural Monument of the Portuguese Republic in 2007.

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